# Factors controlling radiocarbon contents in forest soils

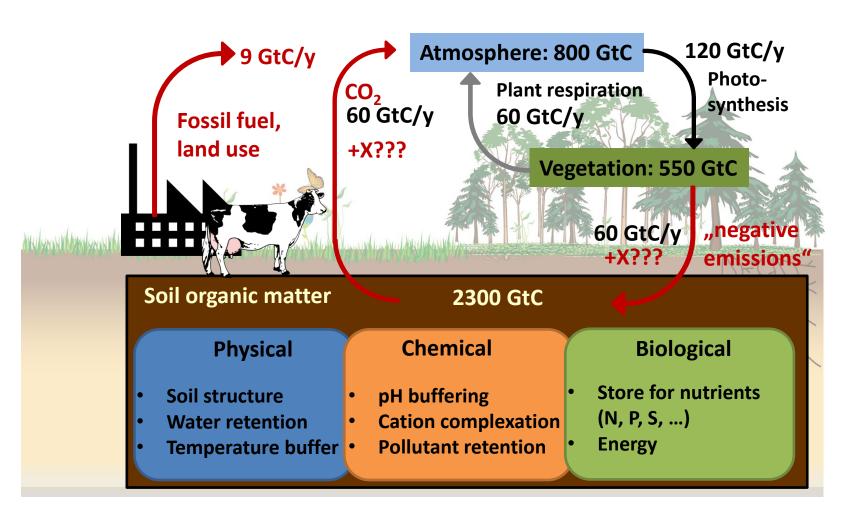
Marion Schrumpf, Ingo Schöning, Ernst-Detlef Schulze, Susan Trumbore

Radiocarbon in the Earth System, September 2017

Max Planck Institute for Biogeochemistry



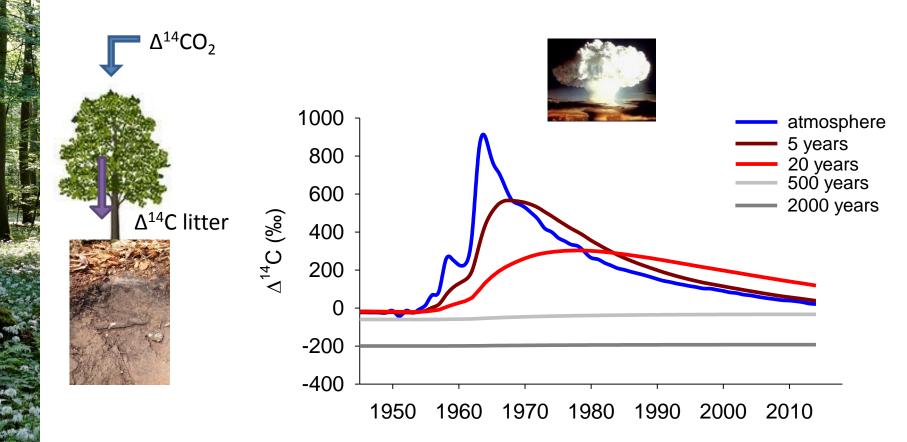
### **Carbon cycling**



What makes soil organic matter persistent and how sensitive is it to global environmental and land use changes?

### Radiocarbon (<sup>14</sup>C) as tracer for soil C dynamics

- Radioactive isotope; half-life: 5730 years; cosmogenic radiation
- Radiocarbon dating: 300 to 50,000 years
- Peak of <sup>14</sup>C in atmospheric CO<sub>2</sub> in 1960s: indicator for C with fast turnover



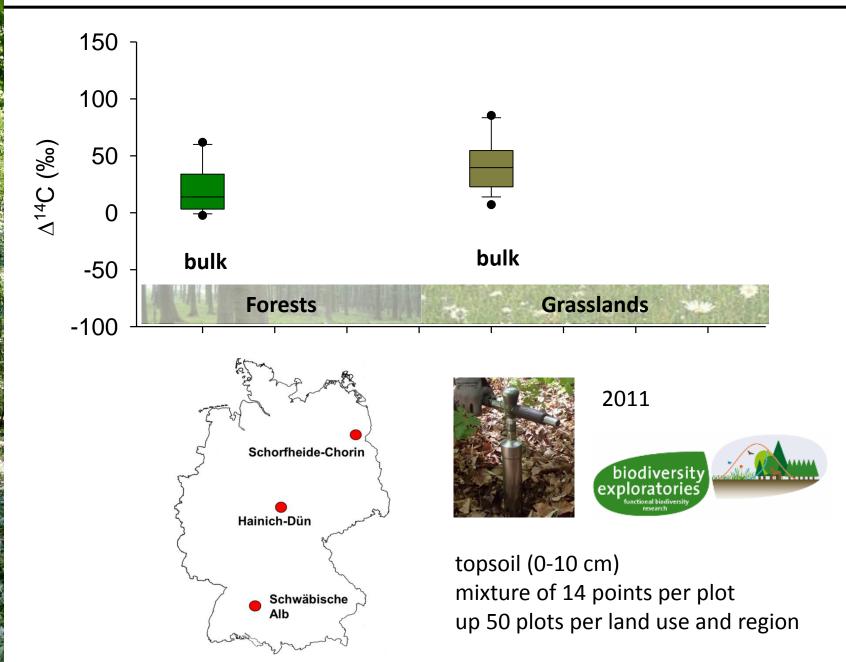
### Outline

- Comparison of <sup>14</sup>C in soils of forests and grasslands: from roots to respired CO<sub>2</sub>
- 2. Application of temporal changes in <sup>14</sup>C
- 3. Studying <sup>14</sup>C along environmental gradients

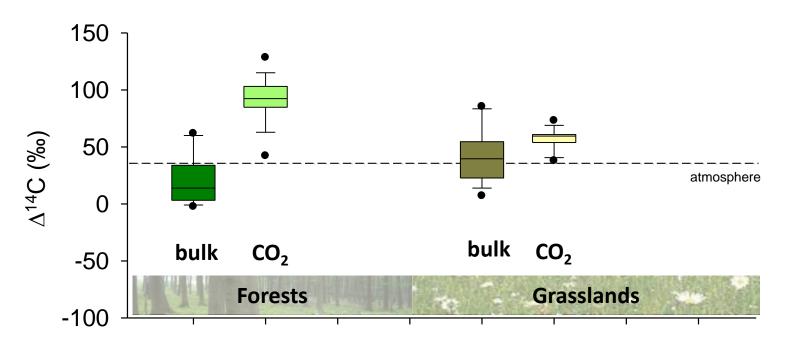
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#### Forests and grasslands: <sup>14</sup>C in bulk soil, 0-10 cm



### Forests and grasslands: <sup>14</sup>C in respired CO<sub>2</sub>



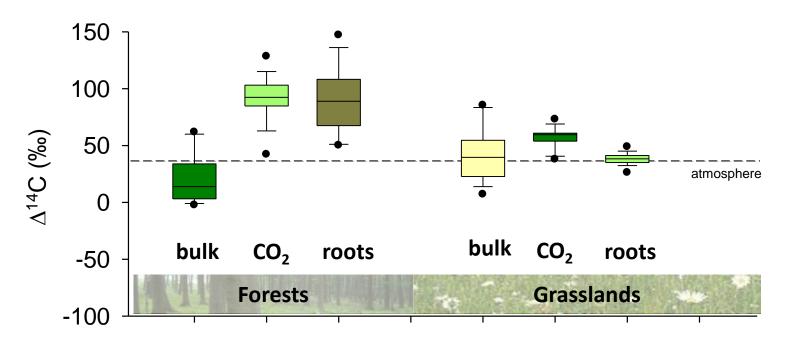


Soil incubation at 20°C and 60% WHC

#### **Respired OC**

- younger than bulk OC
- older in forests than in grasslands
   → slower turnover of fast pool in forests?

### Forests and grasslands: <sup>14</sup>C in living roots





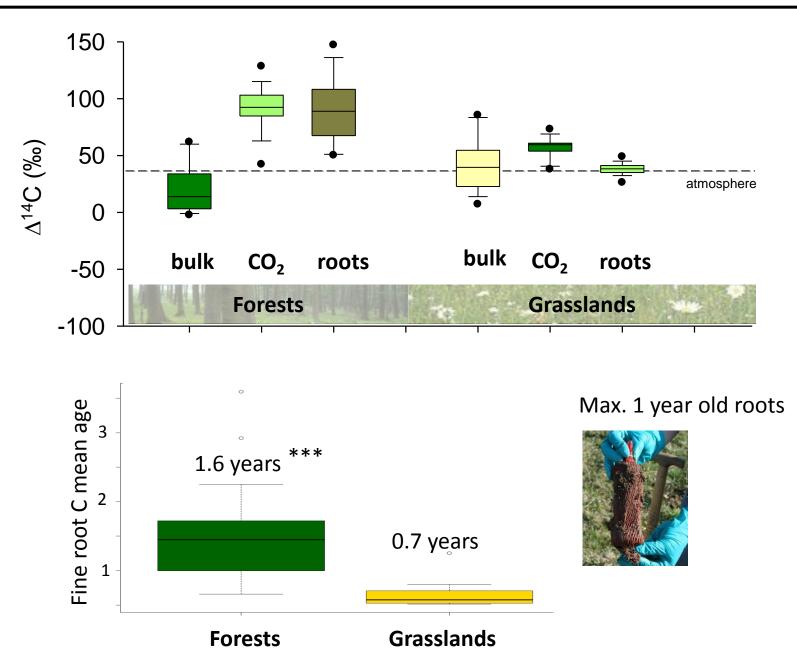
Living root biomass

#### OC in living fine roots

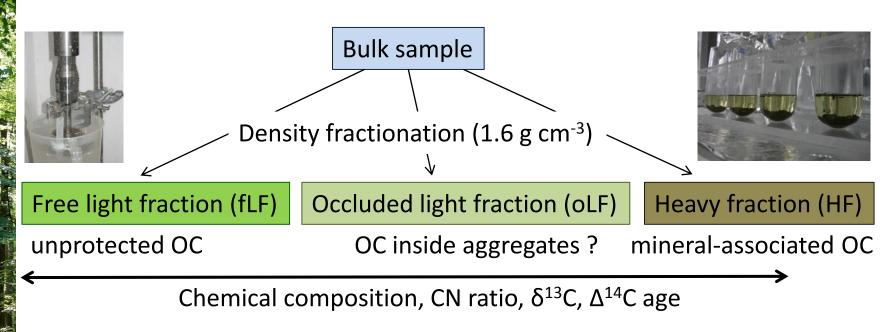
- Older fineroots in forests than in grasslands
   → more longlived roots or usage of storage compunds in forests
- Similar <sup>14</sup>C than in respired CO<sub>2</sub> in forest
   → roots as source for fast soil OC?

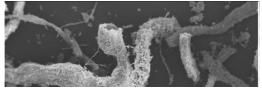
Root data from Solly et al. (2013), Biogeosciences

#### Forests and grasslands: <sup>14</sup>C in living roots

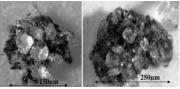


#### Forests and grasslands: <sup>14</sup>C in density fractions

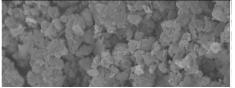




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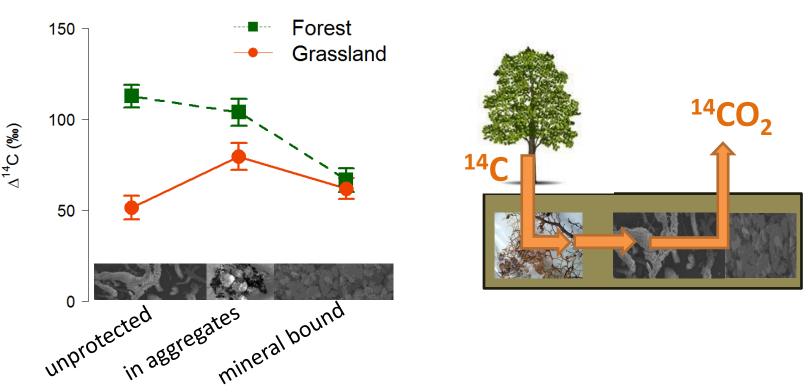
Sarkhot et al 2006 SSSAJ



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#### Forests and grasslands: <sup>14</sup>C in density fractions

Herold et al. (2014), Biogeochemistry



- <sup>14</sup>C from plant litter input can be followed through soil to respired CO<sub>2</sub>
- ightarrow lag time until OC is entering the soil has to be considered
- decomposition in forests really slower than in grasslands?

### Forests and grasslands: root decomposition

Solly et al. (2014)



> 5 000 litter bags at 300 sites



Abiotic conditions

Soil moisture,

Soil temperature

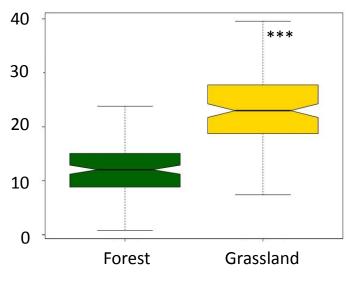
#### Species composition

11

Lignin : N

roots

#### Mass loss (%) after 1 year



#### **Root litter decomposition**

Slower in forests than in grasslands
 → slower soil OC turnover in forests

24

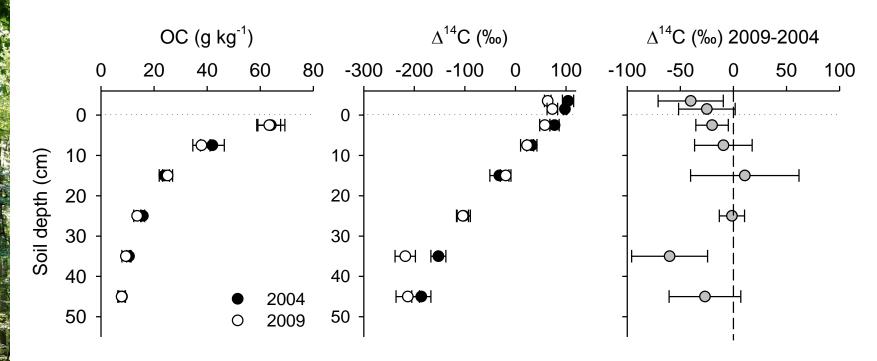
• Driven by climate and litter quality



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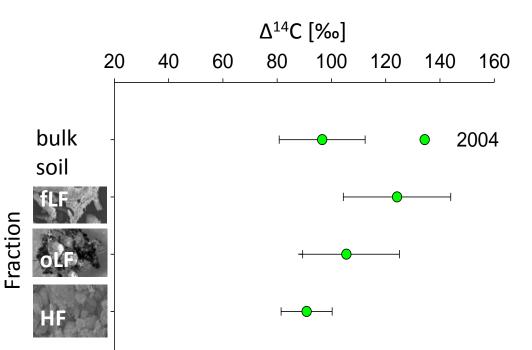
#### Hainich National Park:



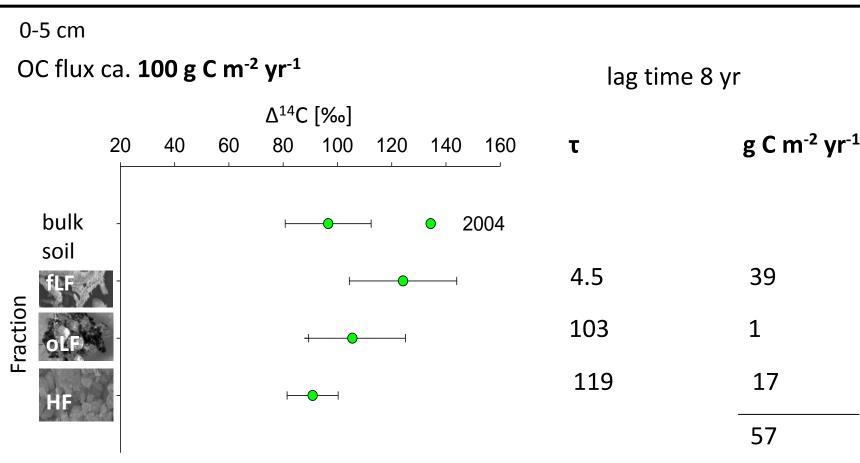
- Decline in <sup>14</sup>C concentrations not detectable below 0-5 cm
   → less contribution of active carbon in subsoils
- We do not know why we find the significant change in 30-40 cm depth...

0-5 cm

#### OC flux ca. 100 g C m<sup>-2</sup> yr<sup>-1</sup>



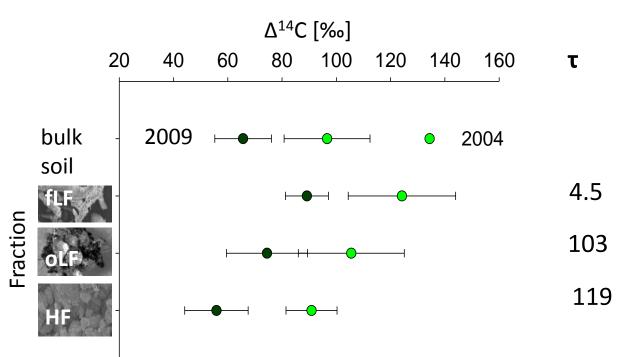
Data: Schrumpf et al. 2015



• Fluxes are too small: an active soil carbon pool is still missing

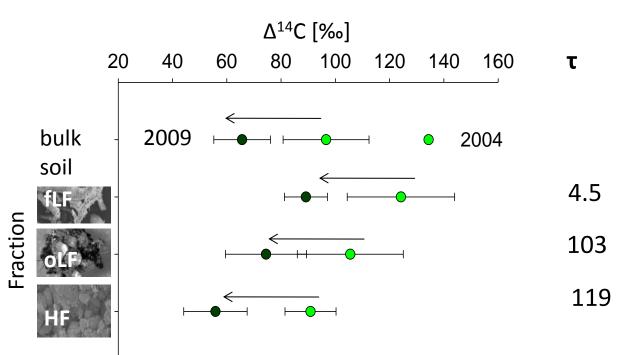
Data: Schrumpf et al. 2015



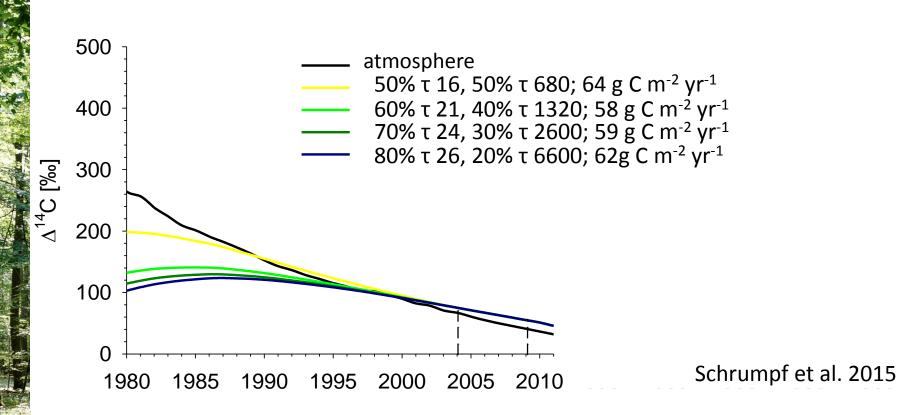


- Given the estimated  $\tau$ , only fLF should respond, but all fractions did:
  - ightarrow also oLF and HF have a portion of active carbon
  - ightarrow both are still composed of mixtures between fast and slow cycling OC

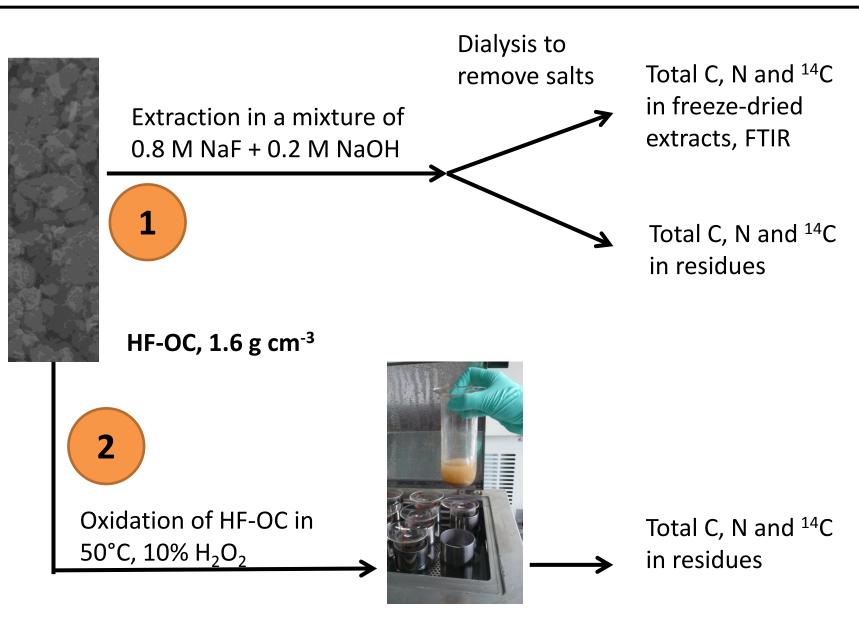
OC flux ca. 100 g C m<sup>-2</sup> yr-<sup>1</sup> (τ: 24 yr)

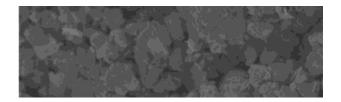


- Given the estimated  $\tau$ , only fLF should respond, but all fractions did:
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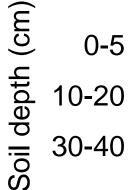
- No unique solution: fast pool: 15-25 years turnover, 50-85% of total OC
- Not much difference between fluxes but: different response to change scenarios total flux: 98-102 g C m<sup>-2</sup> yr<sup>-1</sup>
- Temporal change in <sup>14</sup>C identifies active portions, but not the mechanisms

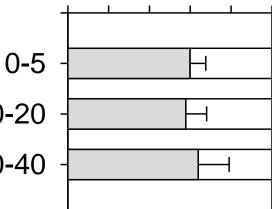




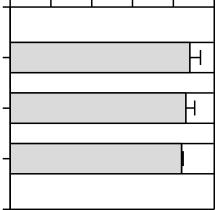
NaF/NaOH extracts

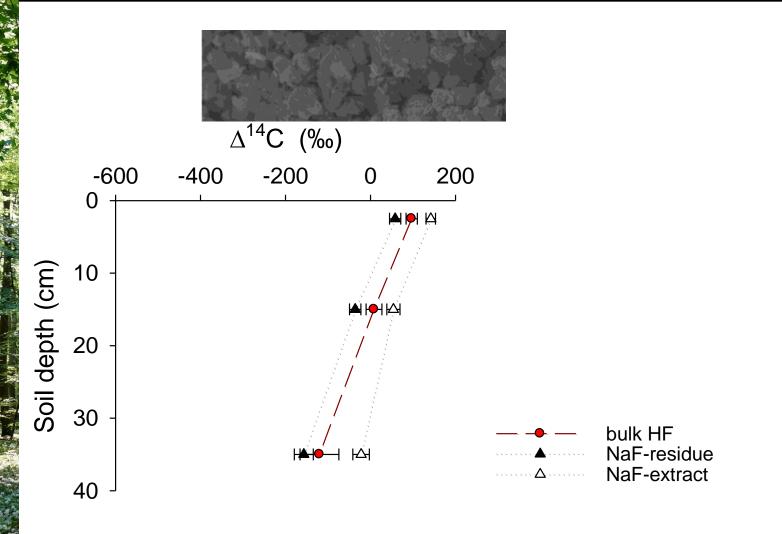
OC extracted (%)



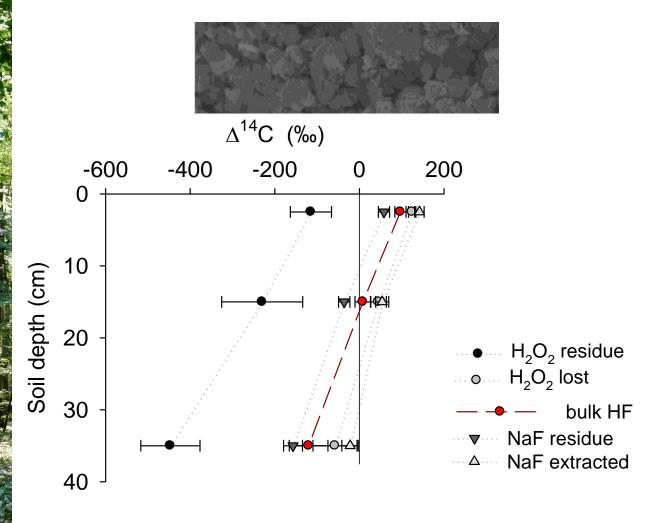


Heated H<sub>2</sub>O<sub>2</sub> treatment OC oxidized (%) 0 20 40 60 80100 0 20 40 60 80100





- OC extracted with NaF is younger than bulk soil HF-OC
- OC extracted with NaF is increasing in age with soil depth



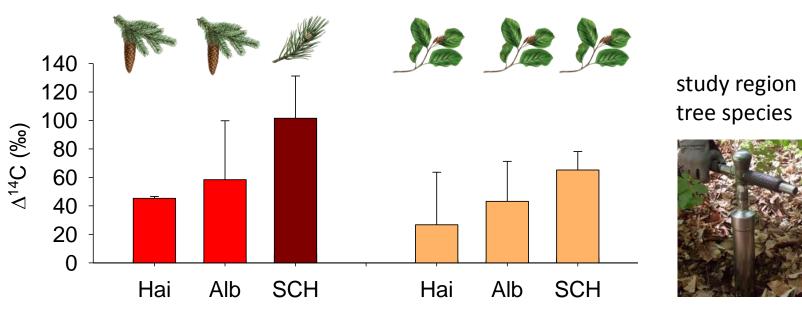
- OC surviving H<sub>2</sub>O<sub>2</sub> treatment is much older than NaF residue
- <sup>14</sup>C of OC lost during H<sub>2</sub>O<sub>2</sub> treatment similar to NaF extracts

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### Tree species: <sup>14</sup>C in mineral soil (0-10 cm)

#### Bulk mineral soil 0-10 cm



<sup>14</sup>C of soil organic matter in mineral soil 0-10 cm:

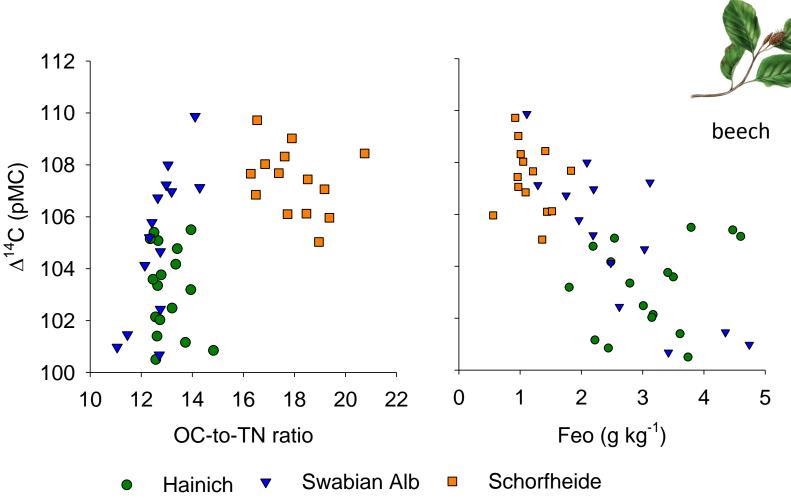
- Larger for coniferous than deciduous species:

   → slower turnover of mineral soil OC for coniferous species
- again: larger difference between regions than between species

 $\rightarrow$  what is driving this huge variation?

### Tree species: <sup>14</sup>C in mineral soil (0-10 cm)

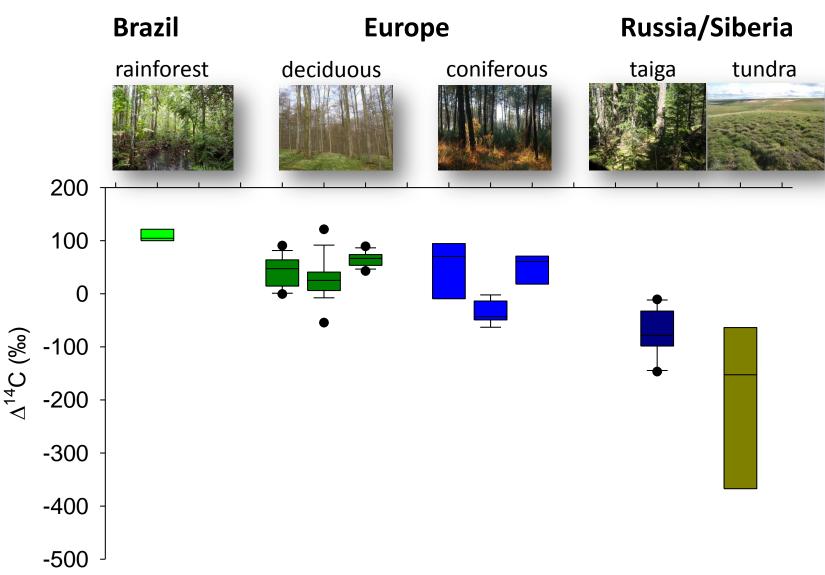




- Study regions differ in climate and various soil properties
- Not possible to identify a single factor explaining <sup>14</sup>C within and across regions

### Comparison of <sup>14</sup>C in soil across large gradients

#### Mineral soil, 0-10 cm, 2004-2013



#### Conclusions

- When using <sup>14</sup>C as proxy for OC turnover consider:
  - soil is mixture of faster and slower cycling fractions
  - lag times between C fixation and the time C enters soil
  - → additional analyses of  $\rm ^{14}CO_2$  or physical / chemical fractions time series
- <sup>14</sup>C values suggest **slower turnover** 
  - for forests than grasslands
  - coniferous than deciduous forests
  - with increasing latitude at the global scale
- <sup>14</sup>C contents in forest soils stronger affected by study regions than by forest type or management

 $\rightarrow$  we do not know the main drivers of <sup>14</sup>C variation (yet)

## Thanks for your attention!



