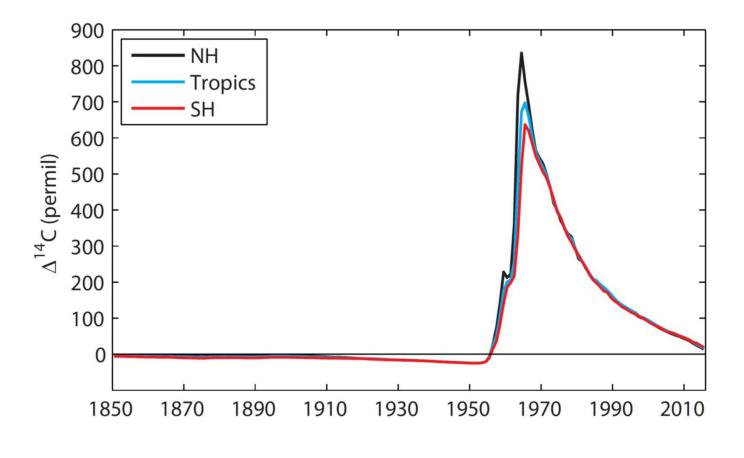
Radiocarbon in Atmospheric CO₂ Heather Graven

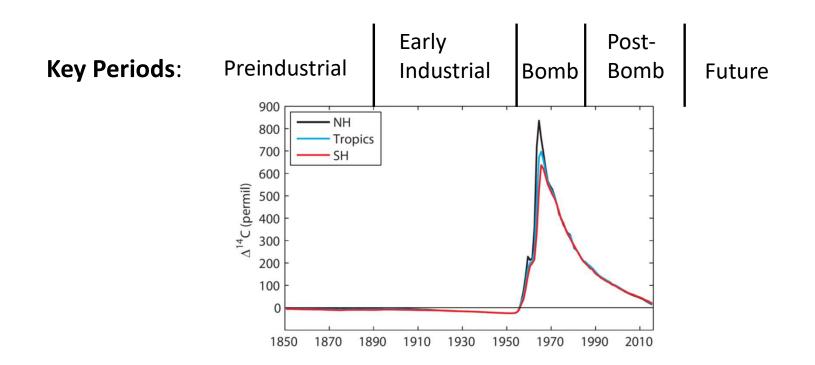


Graven et al. GMDD 2017

Radiocarbon in Atmospheric CO₂

Influences:

Natural ¹⁴C production Fossil fuel emissions Anthropogenic ¹⁴C production Stratosphere-troposphere exchange Ocean-atmosphere exchange Biosphere-atmosphere exchange



Notation and mixing model concept

I will use Δ^{14} C in per mil (‰), including fractionation and age corrections (defined as Δ in Stuiver and Polach 1977)

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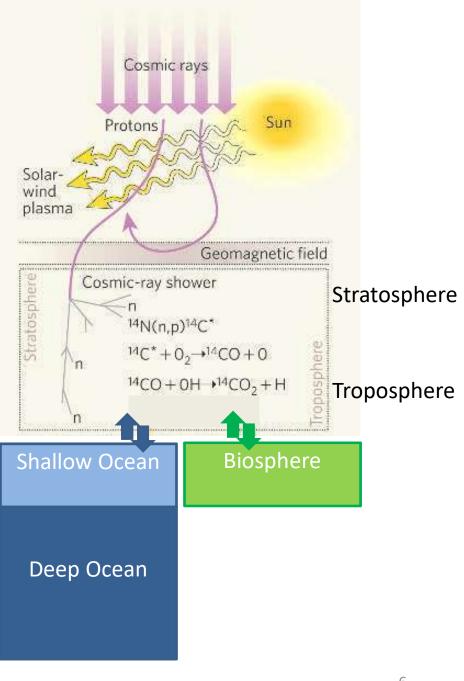
The change in $\Delta^{14}CO_2$ is the difference between the measured (m) and the background (bg) before the source was added:

 $\Delta m - \Delta bg = Cs/Cm (\Delta s - \Delta bg)$

The sign and the magnitude of the change in Δ^{14} C depends on the amount Cs and the difference in Δ^{14} C, the "disequilibrium" ($\Delta s - \Delta bg$)

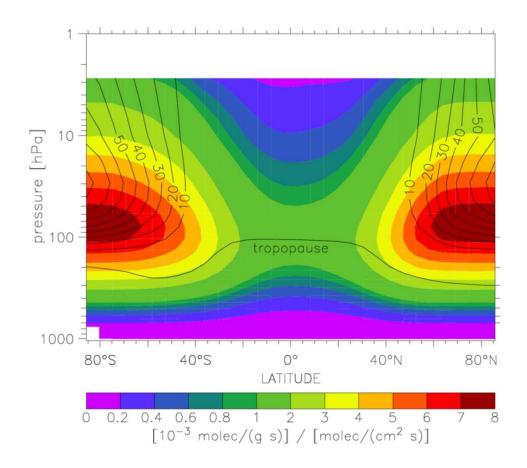
Preindustrial period

- ¹⁴C is produced naturally in the atmosphere
- ¹⁴C participates in the carbon cycle



Jockel and Brenninkmeijer, 2005

Cosmogenic ¹⁴C production occurs primarily in the high latitude stratosphere

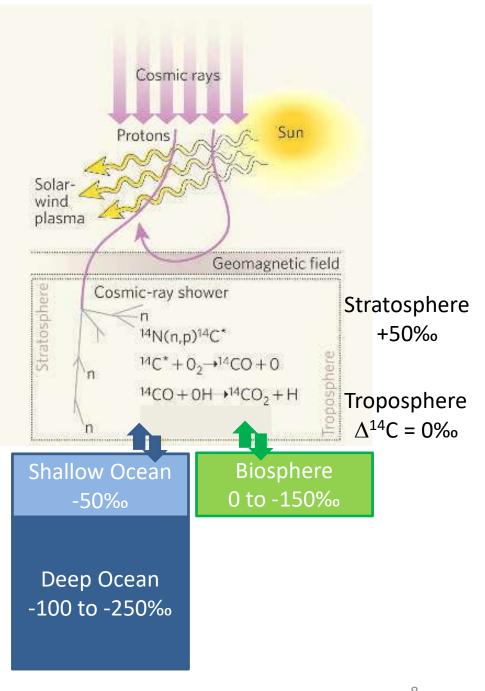


Annual zonal mean galactic cosmic ray induced ^{14}CO production rate (colors), normalized to a global average production rate of 1 molec cm⁻² s⁻¹

From Jöckel et al. 2003, other work by Lingenfelter, Lal, Masarik, Beer, Jockel, Lowe, and new estimate by Lifton et al. 2014

Preindustrial period

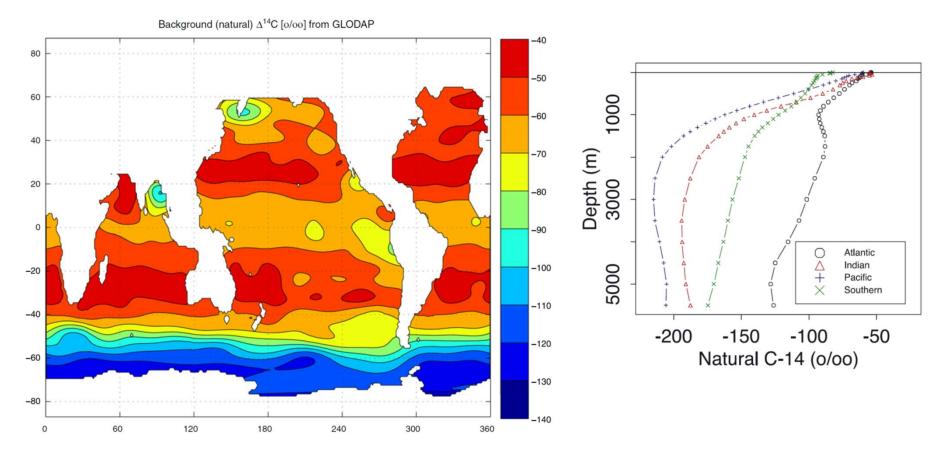
- ¹⁴C is produced naturally in the atmosphere
- ¹⁴C participates in the carbon cycle
- Biospheric and oceanic carbon are depleted in ¹⁴C due to radioactive decay
- Δ^{14} C standard defined to be similar to preindustrial troposphere



Jockel and Brenninkmeijer, 2005

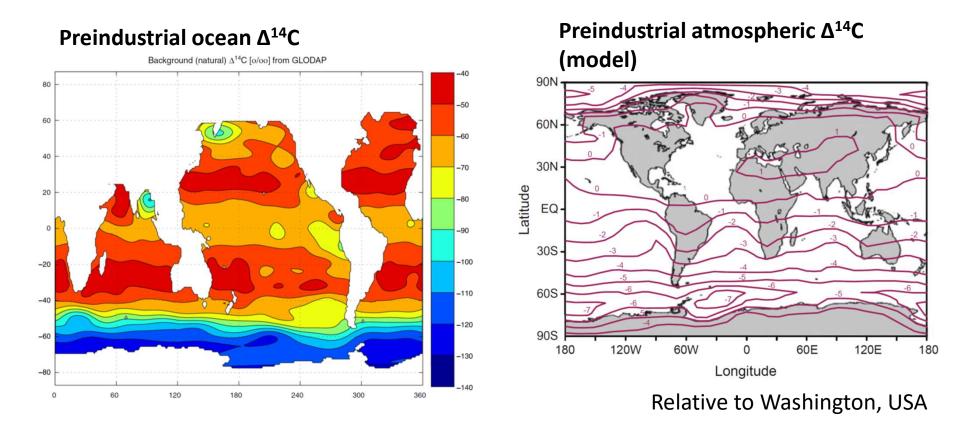
Preindustrial ocean Δ^{14} C distribution

Based on oceanic survey data and estimates of anthropogenic ¹⁴C (GLODAP)



Key et al. 2004; figure from Khatiwala et al. 2012

Preindustrial ocean-atmosphere exchange resulted in higher Δ^{14} C in the Northern Hemisphere, as observed in several studies of tree rings

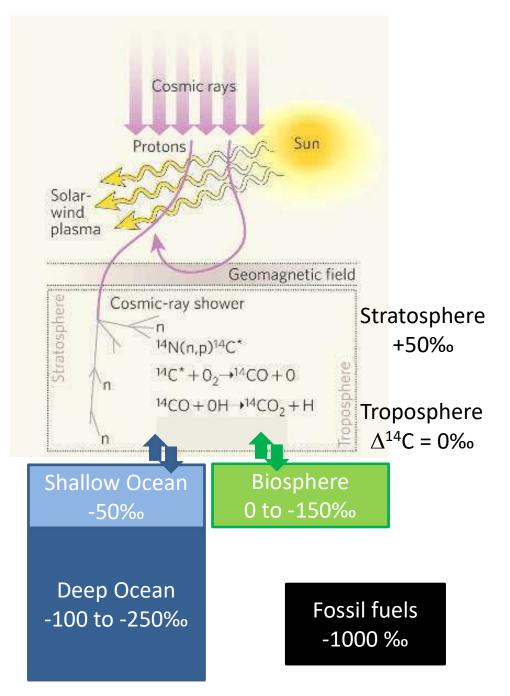


Key et al. 2004; figure from Khatiwala et al. 2012

Braziunas et al. 1995

Early industrial period

- ¹⁴C is produced naturally in the atmosphere
- ¹⁴C participates in the carbon cycle
- Biospheric and oceanic carbon are depleted in ¹⁴C due to radioactive decay
- ∆¹⁴C standard defined to be similar to preindustrial troposphere
- Begin to introduce fossil carbon to the atmosphere through fossil fuel combustion



Jockel and Brenninkmeijer, 2005

"The Suess Effect" – Dilution of atmospheric ¹⁴CO₂ by fossil fuelderived ¹⁴C-free CO₂

Hans Suess, Science, 1955

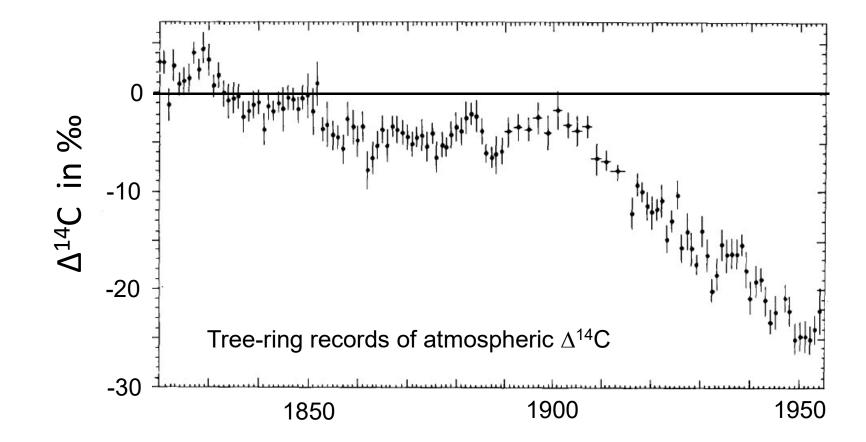
Radiocarbon Concentration in Modern Wood

Redeterminations of the absolute C^{14} concentration in wood carbon and of its variations since the industrial revolution became widespread in the late 19th century were carried out (1) by means of proportional counting of acetylene as described previously (2). No direct determination of the counting efficiency was made, but the counting rates were compared with those from material containing a mass-spectrometrically determined amount of artificial C¹⁴, obtained from the National Bureau of Standards through the courtesy of H. H. Seliger.

Indications of a decrease in the specific C^{14} activity of wood at time of growth during the past 50 yr had been found previously (3). The decrease amounted to about 3.4 percent in two trees from the east coast of the United States. A third tree, from Alaska, investigated at that time, showed a smaller effect. The decrease can be attributed to the introduction of a certain amount of C^{14} -free CO_2 into the atmosphere by artificial coal and oil combustion and to the rate of isotopic exchange between atmospheric

 CO_2 and the bicarbonate dissolved in the oceans.

Early industrial or "Suess" period – "The Suess Effect" in the global carbon cycle



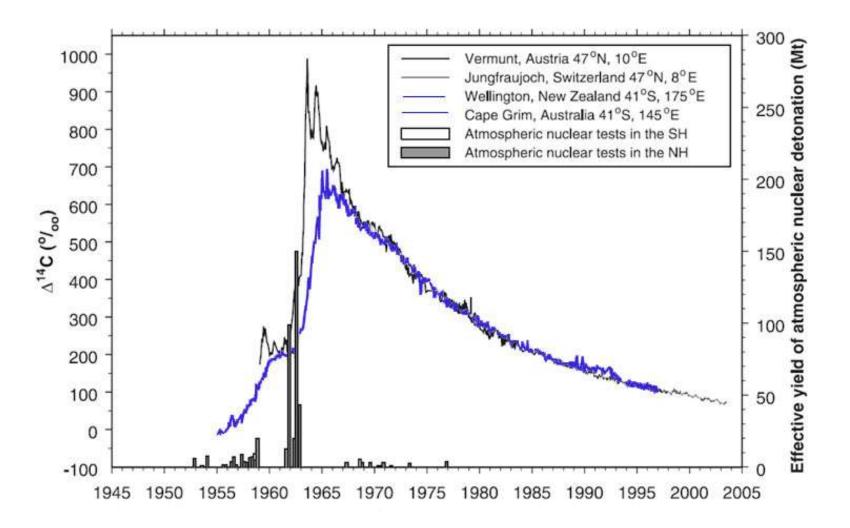
Stuiver and Quay, 1981

"Atom Bomb Effect"—Recent Increase of Carbon-14 Content of the Atmosphere and Biosphere

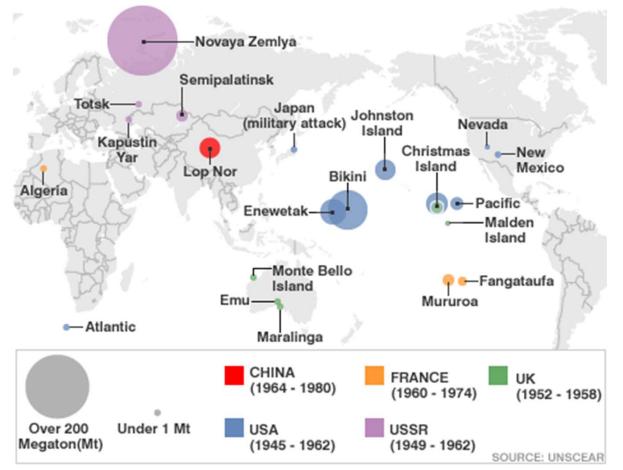
In a study of C¹⁴ variations in nature (1), collection of atmospheric carbon dioxide for the measurement of its C14 activity commenced in New Zealand on 24 Nov. 1954. The first three results were reported (1) in 1955, and the mean value +3.73 percent with respect to the New Zealand wood standard apparently agreed well with Craig's predicted value of +3.68 percent (2). There was, however, in these results an indication that the C14 content of the atmosphere was increasing; hence, the sampling program was continued to check whether or not there was a seasonal variation or a C14 enrichment of the atmosphere by atomic explosions. Nine samples of air have been assayed since 3 June 1955. The results indicate a steady increase in the C14 content of the atmosphere. Duplicate samples collected over the period 12 Dec. 1956 to 23 Mar. 1957 have assayed $+\,7.60\pm0.03$ percent and 8.05 ± 0.30 percent with respect to the New Zealand wood standard. Taking the average value +7.83 percent for the C¹⁴ enrichment of air for the first 3 months of 1957, these results show that the C14 content of the atmosphere of the Southern Hemisphere has increased by 4.10 ± 0.5 percent since February 1955.

Rafter and Fergusson, Science, 1957

"Bomb Period": Nuclear weapons testing caused intense production of ¹⁴C in the atmosphere

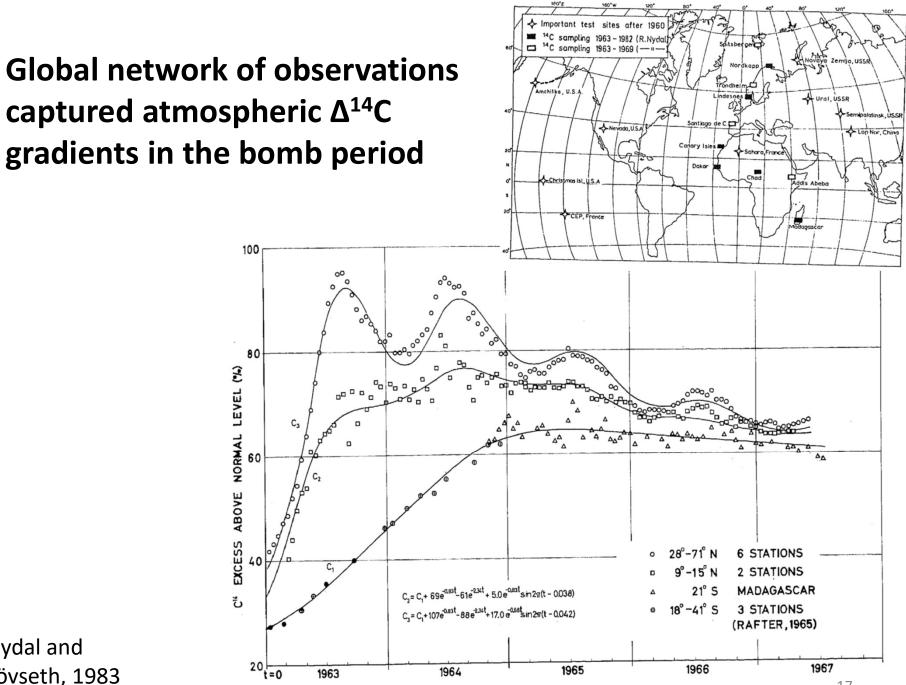


Hua & Barbetti, 2007; Manning et al., 1990; Levin and Kromer, 2004



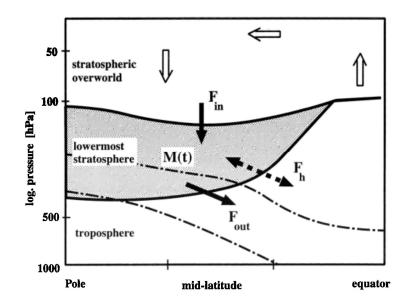
ATMOSPHERIC NUCLEAR TESTS (1945 - 1980)



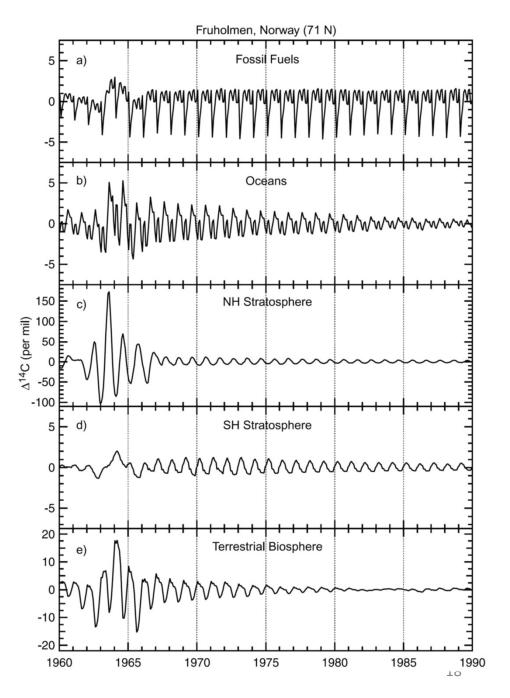


Nydal and Lövseth, 1983 Stratosphere-troposphere exchange caused large Northern Hemisphere seasonal cycles in the early bomb period

Cycles damped slightly by biosphere-atmosphere exchange



Randerson et al. 2002 Appenzeller et al. 1996

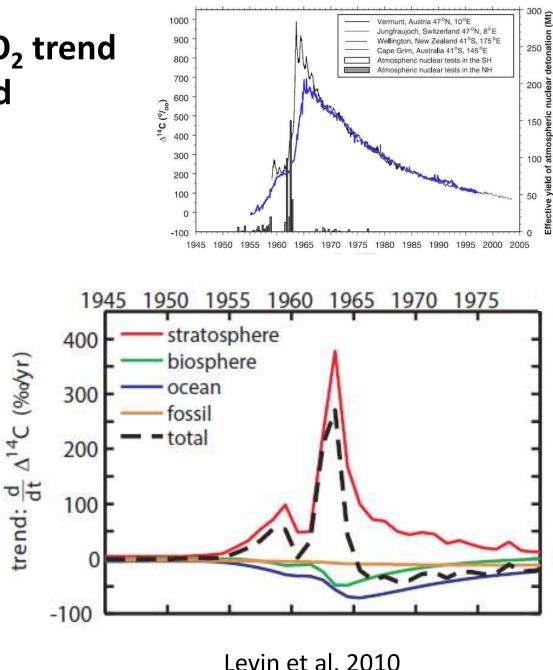


Influences on the $\Delta^{14}CO_2$ trend during the bomb period

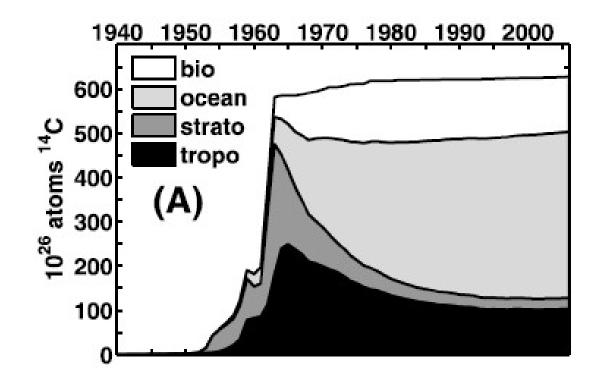
 $\Delta m - \Delta bg = Cs/Cm (\Delta s - \Delta bg)$

Ocean and biosphere exchange with the atmosphere introduced carbon with lower Δ^{14} C into the atmosphere

Ocean and biosphere exchanges counteracted and then overpowered the stratospheric influence after the majority of testing ceased



Inventory of bomb ¹⁴C in different reservoirs evolved over time as a result of carbon cycling



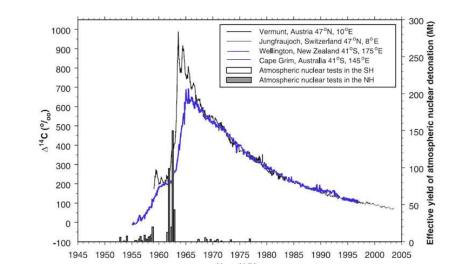
Estimates of total 14C produced by bomb strength and number

Some ¹⁴C is continually emitted from nuclear power plants (about 10% of natural production rate)

Ocean and atmosphere bomb ¹⁴C inventories constrained by observations

Biospheric inventory by residual

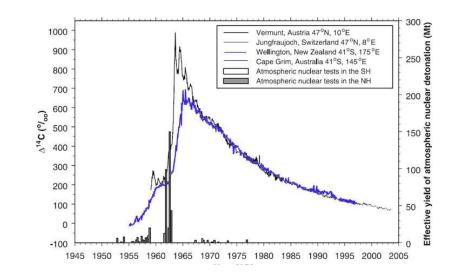
Naegler and Levin 2009



What do you think the atmospheric Δ^{14} C history would have looked like if:

Discussion

the exchange between the stratosphere and the troposphere was much slower?

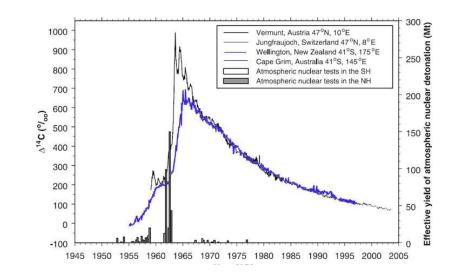


What do you think the atmospheric Δ^{14} C history would have looked like if:

Discussion

the exchange between the stratosphere and the troposphere was much slower?

the exchange between the atmosphere and the ocean was much faster?



What do you think the atmospheric Δ^{14} C history would have looked like if:

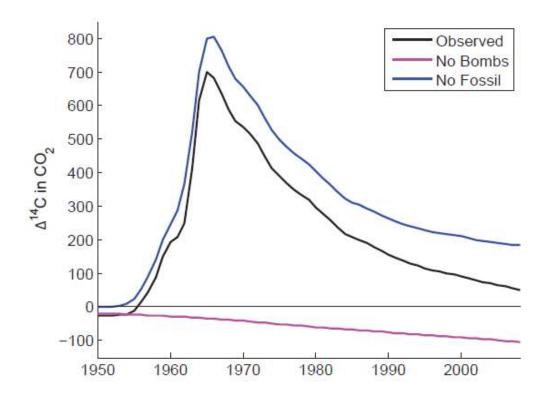
Discussion

the exchange between the stratosphere and the troposphere was much slower?

the exchange between the atmosphere and the ocean was much faster?

there were no fossil fuel emissions, but there were nuclear weapons tests?

Imagine if there were no fossil fuel emissions, but there were nuclear weapons tests - what do you think the atmospheric Δ^{14} C history would have looked like?



Simulations using simple carbon cycle model

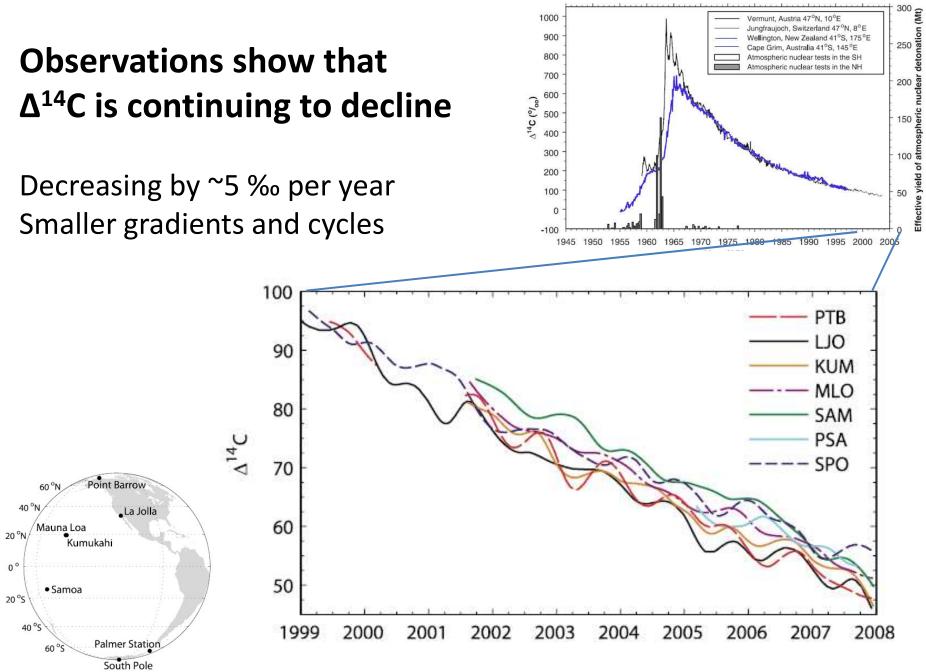
Scenario w/o fossil fuel emissions:

- Similar to observed trend through 1980s
- Weaker trend than observed after 1990

Scenario w/o bomb testing:

 Indicates atmospheric ∆¹⁴C would now have been -100 ‰ due to fossil fuel emissions

Graven, unpublished



Graven et al. 2012 ²⁵

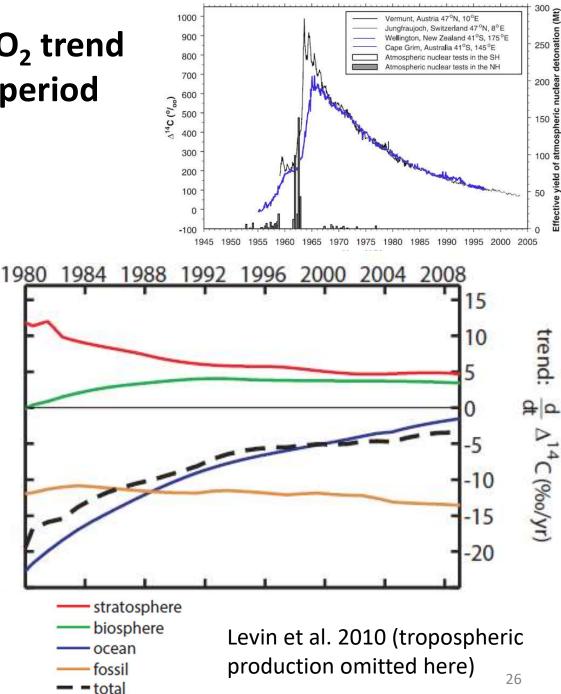
Influences on the $\Delta^{14}CO_2$ trend during the post-bomb period

 $\Delta m - \Delta bg = Cs/Cm (\Delta s - \Delta bg)$

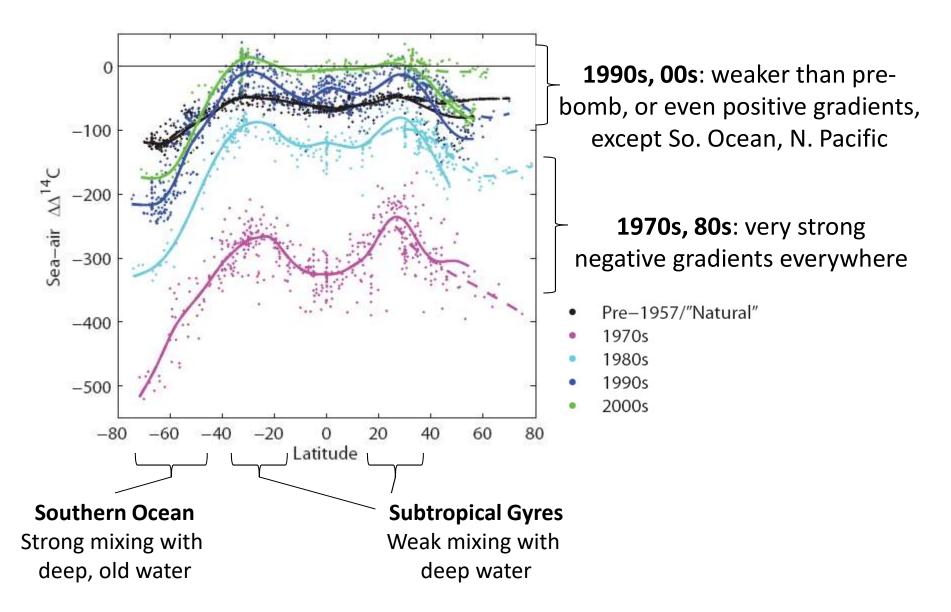
Stratosphere and biosphere exchange have recently had relatively steady positive influences

Ocean exchange strongly weakened

Fossil fuel emissions are now the largest influence on the $\Delta^{14}CO_2$ trend

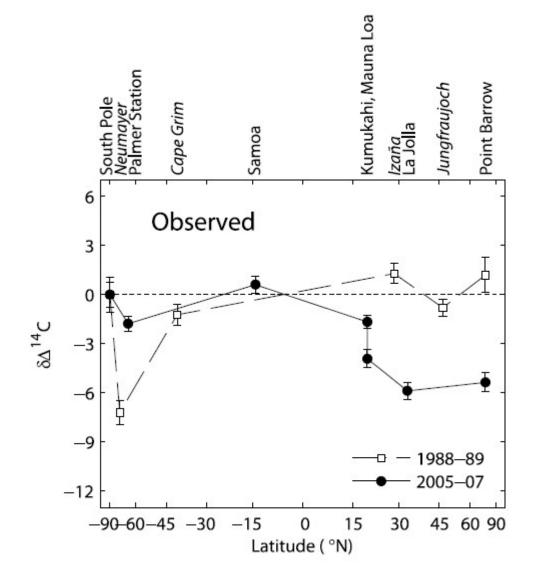


Observed sea-air Δ^{14} C gradients show switch to 14 C outgassing



Linick 1978; Fairhall 1985; Broecker 1985; Nydal 1998; Key 2004 and many others

Observed recent shift in North-South gradient



 Δ^{14} C relative to the South Pole in 1988-89 and 2005-07

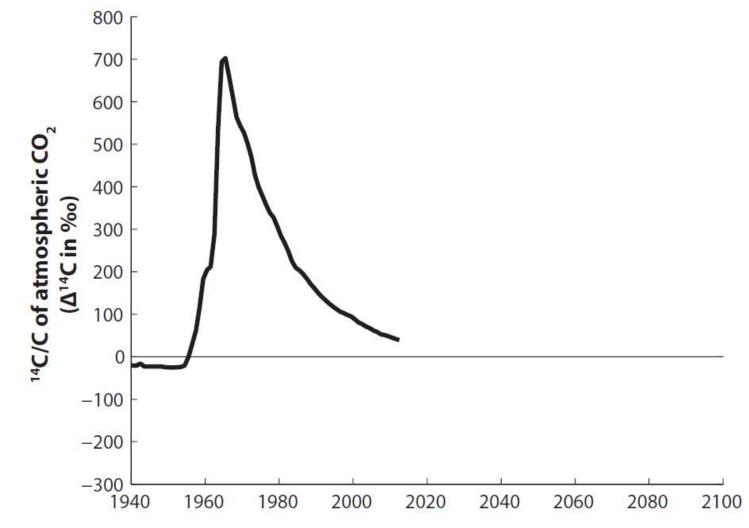
Main influences on gradient:

- Fossil fuel emissions in the North
- Ocean exchange in the South

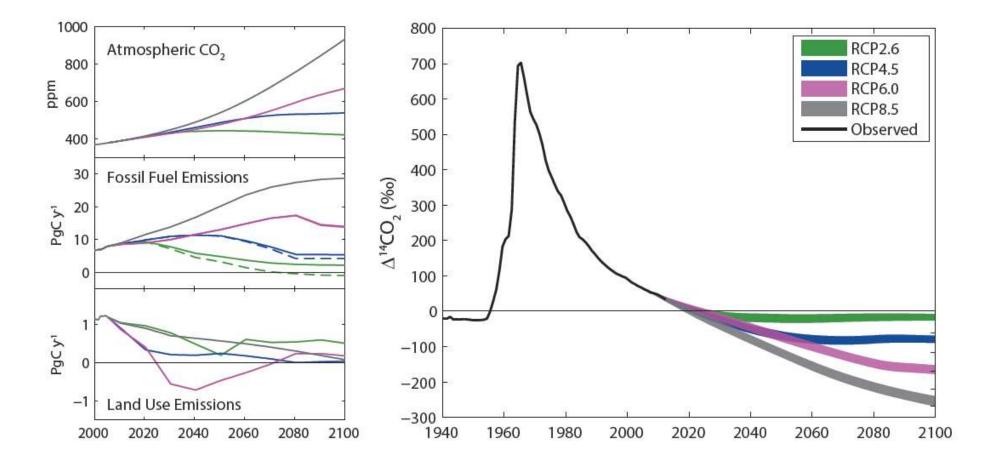
Shift indicates change in relative strength of fossil fuel and ocean influences

Gradient is now opposite to preindustrial gradient

How will atmospheric Δ^{14} C evolve over this century?



A wide range of $\Delta^{14}CO_2$ trajectories are simulated using the RCPs and a simple carbon cycle model



Range of emission scenarios – Representative Concentration Pathways (RCPs) from IPCC AR5

Graven, PNAS, 2015₃₀

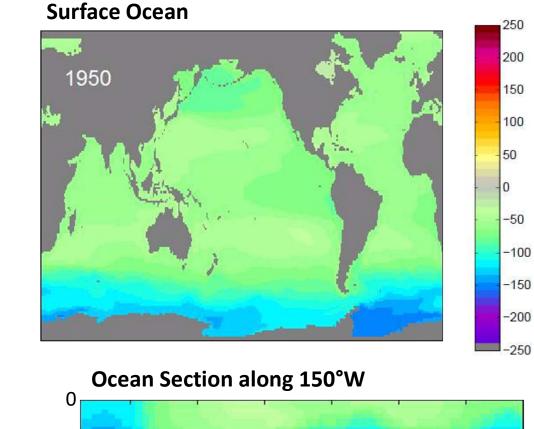
Model simulations of future ocean $\Delta^{14}C$

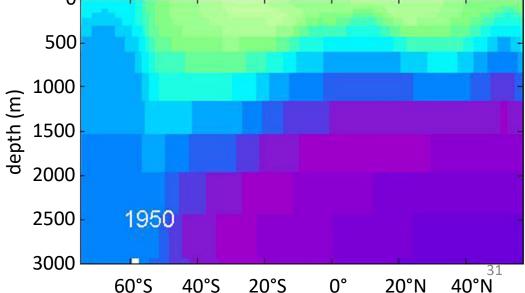
ECCO ocean model forced with atmospheric Δ^{14} C observations and projections (Khatiwala 2007; Graven et al. 2012)

Constant ocean circulation

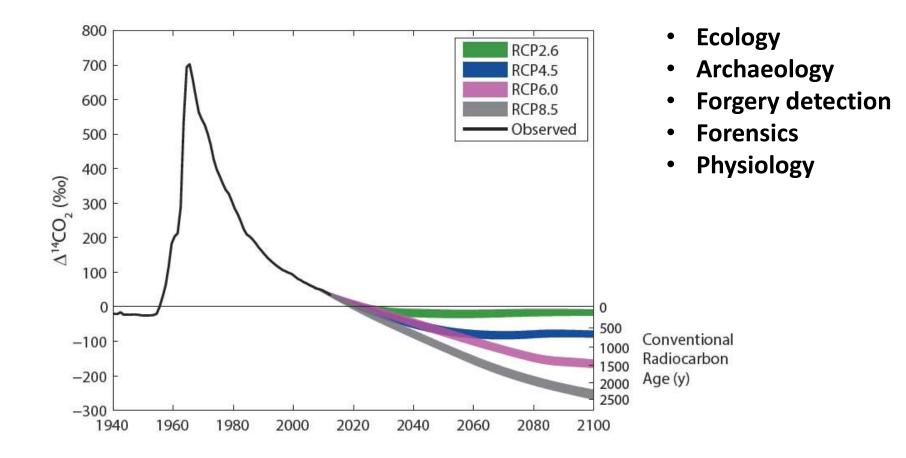
Gas exchange velocity of 15 cm/hr (Sweeney et al. 2007)

Khatiwala et al., in rev.





Forthcoming changes in $\Delta^{14}CO_2$ have implications for various applications of ^{14}C



Graven, PNAS, 2015₃₂

Summary

Nuclear and fossil fuel perturbations have caused dramatic changes to Δ^{14} C in atmospheric CO₂

Disequilibria with the stratosphere, ocean and biosphere have evolved over time and, in some cases, changed sign

Now that several decades have past since the bomb tests, fossil fuel emissions are becoming more and more dominant in driving Δ^{14} C trends and gradients

 Δ^{14} C in atmospheric CO₂ will reach 0 ‰ in the next couple of years, and depending on fossil fuel emissions, may decline to -250 ‰ by the end of the century, impacting various ¹⁴C applications

Carbon isotopes in CMIP6

Motivation:

- Improved understanding of carbon cycle processes
- Model evaluation

We are providing historical and future forcing data for atmospheric $\delta^{13}CO_2$ and $\Delta^{14}CO_2$

Data available at https://esgf-node.llnl.gov/search/input4mips/ Graven et al. GMDD 2017

Ocean-only simulations: OMIP (Ocean MIP) Historical ocean-only simulations with observed climate Orr et al. GMD 2017

Land and Ocean simulations:

C4MIP (Coupled Climate-Carbon Cycle MIP) Historical and future simulations with coupled climate Jones et al. GMD 2016

